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ENSO Impact on Winter Precipitation in the Southeast United States through a Synoptic Climate Approach

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Abstract: The ENSO impact on winter precipitation in the Southeast United States has been analyzed from the perspective of daily weather types (WTs). We calculated the dynamic contribution associated with the change of frequency of the WTs and the thermodynamic contribution due to change of the spatial patterns of the environmental fields of the WTs. Six WTs have been obtained using a k-means clustering analysis of 850 hPa winds in a reanalysis data from November to February of 1948–2022. All the WTs can only persist for a few days. The most frequent winter weather type is WT1 (shallow trough in eastern U.S.), which can persist or likely transfer to WT4 (Mississippi River Valley ridge). WT1 becomes less frequent in El Niño years, while the frequency of WT4 does not change much. WTs 2–6 correspond to a loop of eastward propagating waves with troughs and ridges in the mid-latitude westerlies. Three WTs with a deep trough in the Southeast U.S., which are WT2 (east coast trough), WT3 (off east coast trough) and WT6 (plains trough), become more frequent in El Niño years. The more frequent deep troughs (WTs 2, 3 and 6) and less frequent shallow trough (WT1) result in above-normal precipitation in coastal Southeast U.S. in the winter of El Niño years. WT5 (off coast Carolina High), with maximum precipitation extending from Mississippi Valley to the Great Lakes, becomes less frequent in El Niño years, which corresponds to the below normal precipitation from the Great Lakes to Upper Mississippi and Ohio River Valley in El Niño years, and vice versa in La Niña years. The relative contribution of the thermodynamic and dynamic contribution is location dependent. In the east coast, the two contributions are similar in magnitude.

Keywords: El Niño; Winter Weather Types

Citation: Qian, J.-H.; Viner, B.; Noble, S.; Werth, D.; Li, C. **Winter weather types and ENSO impact on precipitation in the Southeast United States.** *Atmosphere*, **2022**, *04*, x. <https://doi.org/10.3390/xxxxx>

Academic Editor: First name Last-name

Received: date

Accepted: date

Published: date

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1. Introduction

The influence of the El Niño – Southern Oscillation (ENSO) on U.S. climate is strongest during the winter [1, 2]. The problem is usually studied based on seasonal means, such as the ENSO impact on the mean precipitation in December–February [3]. As we experience in everyday life, climate and its variability are actually manifested by daily weather. Therefore, it is equally important to learn how the ENSO impacts the daily weather. The problem will be examined and quantified using a weather typing (WT) analysis in Southeast United States (SEUS) for the winter in the current study. Similar method has been applied in the study of the impact of the ENSO in other places of the world such as in the Northeast U.S. [2, 4] and in the Southeast Asia [5, 6, 7, 8] as well as on spring precipitation in the SEUS [9]. The rationale of the WT analysis and reference therein can be found in Qian et al. [9].

From the large-scale point of view, winter WTs in the SEUS are closely related to the North Atlantic Subtropical High (NASH) south of the Gulf coast (along about 30 °N) and westerly winds north of 30 °N, as shown in the monthly climatology in Fig. 1. However, the monthly precipitation and 850 hPa circulation of the four months from November to February look almost the same, especially in the deep winter months of December to February with the anticyclonic circulation of the NASH withdrawn further to the

south. Winter rainfall is often produced by southwesterly flows ahead of a cold front that brings moisture from the warmer Gulf of Mexico. The location of precipitation should depend on the weather regime of the day, especially the longitude of the trough. However, the ridges and troughs cannot be seen from the monthly or seasonal mean figures. As will be seen shortly, the weather typing analysis brings up the dynamic view of these propagating ridges and troughs. Cold air damming east of the Appalachian Mountains can also bring cold weather to the south in the winter [10, 11]. Geographically special for the east coast, the strong temperature gradient between the warm Gulf Stream off the east coast and the cold coastal land in the cold seasons also forms a baroclinic belt, which corresponds to the quasi-stational east coast trough and storm track, as seen from the heavy precipitation along the east coast in Fig. 1. All these are important factors that affect winter weather in the SEUS.

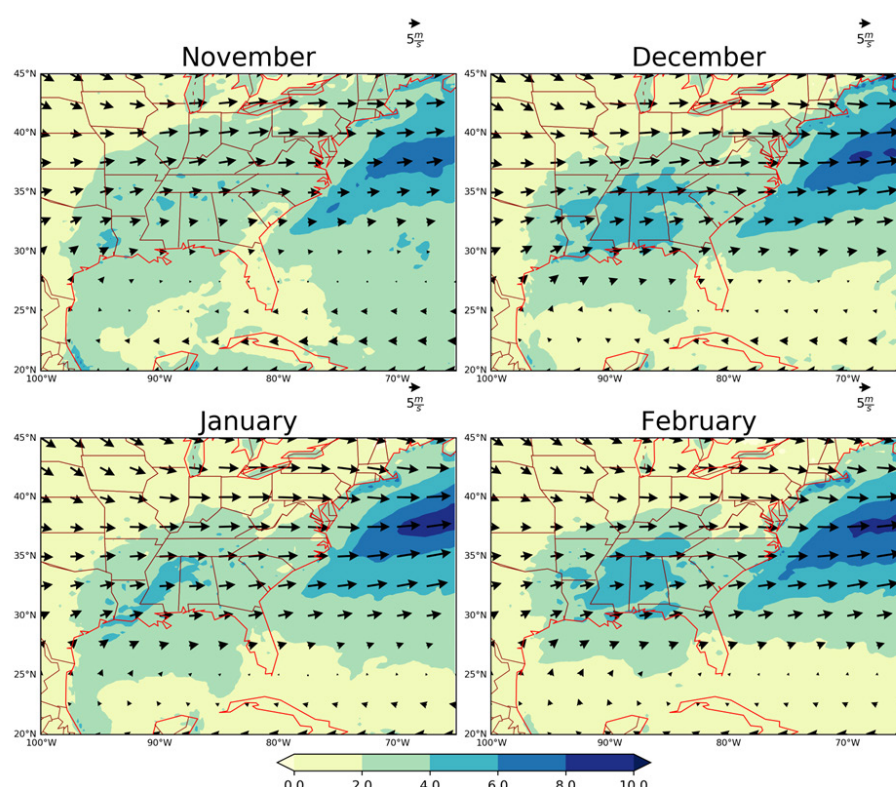


Figure 1. Monthly climatology of precipitation (shades, mm/day) and 850 hPa horizontal circulation (arrows, m/s) in the winter in the Southeast U.S. The mean winds are based on the NRP reanalysis of 1948–2022. The mean precipitation is based on the CMORPH (NOAA Climate Prediction Center–CPC Morphing Technique) satellite estimated data in 1998–2020.

In the following, we will first describe the data and method briefly, then examine the spatial pattern and frequency of the winds and precipitation of the WTs. Finally, we will use the characteristics of the WTs to interpret the ENSO impact on precipitation in the SEUS.

2. Data and Weather Typing Analysis

2.1. Data

As in Qian et al. [7, 9], the 850 hPa wind components of the NCEP–NCAR Reanalysis Project–I (NRP) data [12], starting from January 1, 1948 to February 28, 2022, are used for the WT analysis. To examine the precipitation pattern of each of the WTs, the satellite estimated 3-hourly precipitation data from CMORPH (NOAA Climate Prediction Center–CPC Morphing Technique, $0.25^\circ \times 0.25^\circ$, 1998–2020, covering both land and sea) [13] are used to plot the precipitation maps. The CPC unified gauge-based analysis of daily

precipitation over the contiguous United States (CONUS) at quarter-degree grids in 1948–2022 (<https://psl.noaa.gov/data/gridded/data.unified.daily.conus.html>, covering only land area) is used for analyzing the ENSO impact on precipitation.

2.2. WT Analysis

The domain used for the WT analysis, 100°W–65°W and 20°N–45°N, is the same as was used to analyze the WTs in the warm season in the SEUS [9]. In this paper, we analyze November to February. We also extend the years of analysis by including all currently available NRP data: 1948–2022.

As in Roller et al. [2] and Qian et al. [7, 9], the classifiability index (CI) is calculated to determine the optimum number of clusters for the dataset, with large CI indicating a better separation between clusters. The optimum number of clusters (k) corresponds to the peak in CI following the initial decrease. The daily weather in the SEUS in the winter can be best represented by 6 clusters ($k=6$) (Fig. 2).

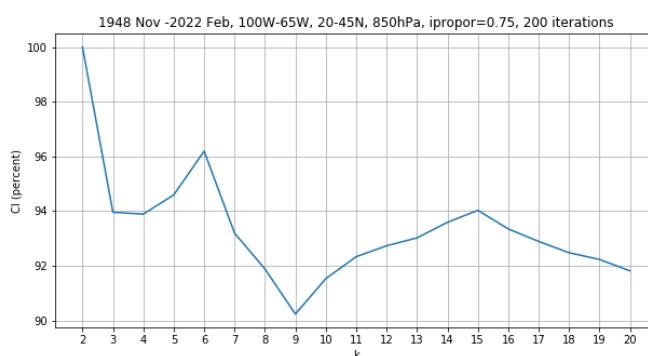


Figure 2. The classifiability index (CI) of the k-means clustering analysis using 850 hPa zonal and meridional wind component (u and v) in the Southeast U.S. and the surrounding region (100°W–65°W, 20°N–45°N) in November–February, 1948–2022; the optimum number for the k-means analysis is $k=6$.

3. Results

3.1. Winter weather types in the SEUS

The spatial distribution of the 850 hPa winds for the six WTs, along with the associated mean precipitation of each, is shown in Fig. 3 and the brief description is given in Table 1. WT1 represents a broad and shallow trough over central-east US. WT2 is an east-coast trough with precipitation ahead of the trough along the east coast from Florida to the mid-Atlantic coast, but with the maximum precipitation over the costal ocean along the Gulf stream. Behind the trough, northwesterly dry flow takes place over the Midwest and Great Plains. The Gulf coast is also quite dry. WT3 is an off-east coast trough with maximum precipitation over the ocean, while the land in central and eastern U.S. is dry. WT4 features a ridge over the Mississippi River Valley, with a trough and rainy area in the Atlantic Ocean quite far away from the east coast. The east coast and the Great Lakes area are dry. The southern Great Plains get some rainfall with the moisture supplied by the southerly flow from the Gulf coast. In WT5, a trough is over the western Great Plains, with the southwesterly flow brings forth moisture and precipitate over the Mississippi River Valley. An anticyclone is centered off the east coast of Carolinas. The east coast is quite dry. In WT6, a deep trough is over the eastern Great Plains and the Mississippi and Missouri river valleys. The strong southwesterly flow east of the trough transports moisture from the Gulf of Mexico and dumps heavy precipitation in the Mississippi River Valley and SEUS.

Comparing the winter WTs to those WTs in March–October given in [9], some WTs are similar to or even identical to each other. For instance, the WT2 in Fig. 3 here is quite

similar to the WT1 (ECT) in [9]. WT4 here in Fig. 3 is almost identical to the WT2 (MRVR) in Qian et al. [9]. While WH6 here in Fig. 3 is similar to the WT3 (PT) in Qian et al. [9]. It is reasonable that some WTs in November-February may share similar features of the WTs in the transitioning seasons of spring and fall contained in the period of March-October in Qian et al. [9].

Weather Type	Description
WT1	Shallow trough (ST) over central-east US
WT2	East coast trough (ECT)
WT3	Off-east-coast trough (OECT)
WT4	Mississippi River Valley Ridge (MRVR)
WT5	Western plains trough (WPT)/Off coast Carolina high (OCCH)
WT6	Plains trough (PT)

Table 1. Six weather types in the winter and their 850 hPa flow patterns.

WT5 bears some similarity to the WT6 (Fig. 3), but with a shallower trough in western Great Plains. WT5 is also somewhat similar to the warm season WT6 (Carolina High, [8]), but the anticyclone center in the WT5 here is off the Carolina coast, while the warm season WT6 has the Carolina High centered on the coast.

We cannot find similar warm season WTs similar to the winter WT1 and WT3, therefore, they are distinctively winter WTs. In the winter, with the contrast of cold continent and warm ocean, the general circulation favors a quasi-stationary trough along the east coast of the continents. This is manifested by the occurrence of shallow trough in coastal land (WT1) and a deep off-shore trough (WT3).

It is also worth to note that precipitation is heavy over the Great Lakes (near the northern border of the domain in Fig. 3), as compared to the surrounding lake shore region, in WTs 5 and 6, in which southwesterly flow brings moisture toward the Great Lake region. In the winter, surface temperature over the lakes should be warmer than that over the land. Therefore, moist air and the warm lake surface both work to increase the conditional instability of the atmosphere thus enhance precipitation over the lakes.

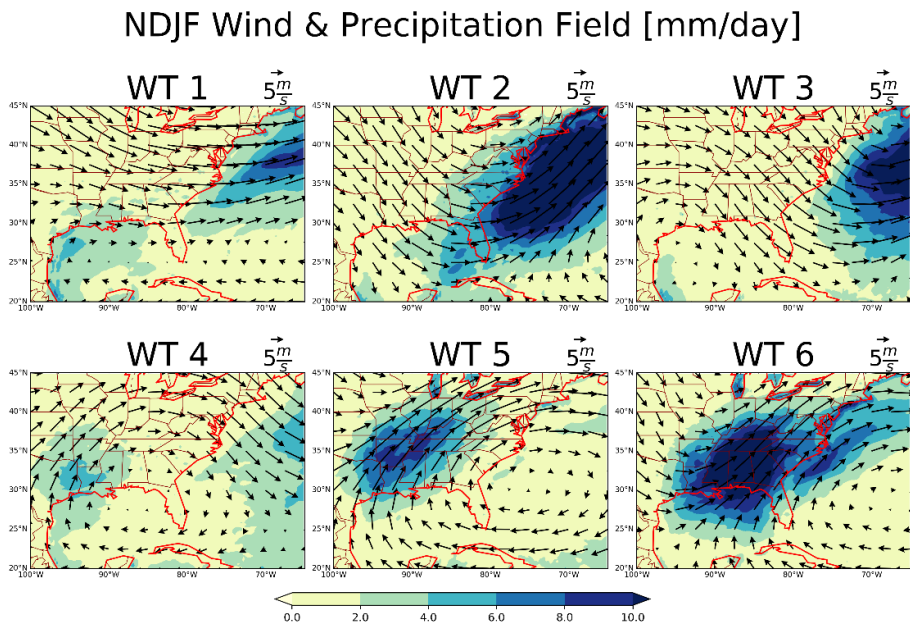


Figure 3. The average precipitation (shade, mm/day) and 850 hPa winds (arrows) of the six WTs. WTs 1-6 can be briefly described as Shallow Trough (ST), East Coast Trough (ECT), Off East Coast Trough (OECT), Mississippi River Valley Ridge (MRVR), Western Plains Trough (WPT), and Plains Trough (PT), respectively.

3.2. Frequency of the WTs

During November 1 – February 28 of 1948–2022 (120 days per year for 74 years, 8880 days in total), the percentage of the total number of days with WT 1–6 are 21.3%, 11.7%, 13.9%, 19.0%, 17.5% and 16.7%, respectively. The occurrence of the daily WTs through the season is shown in Fig. 4. All six WTs sporadically occur in the winter season, as can be seen from the top panel of Fig. 4. WT1, a distinctive winter WT, occurs quite frequent in every month of the winter, especially in December and January; therefore, it is the most frequent WT in the center of the winter. In the winter, westerly winds are the strongest in the mid-latitude due to the strongest baroclinicity (Fig. 1), which corresponds to the quite zonal flow of WT1. WT2 (ECT), the least frequent winter WT, occurs slightly more in January and early February. As mentioned before, the WT2 (ECT) here is similar to the WT1 (also ECT) in [9] that frequently occurs in March–April. So ECT is common in the winter and spring in SEUS. WT3 (OECT), the second least frequent winter WT, occurs more in January and February. WT4 (MRVR, the second most frequent WT) occurs more frequent in November (especially in the early part of the month), and least frequent in January. Precipitation map of WT4 indicates that it corresponds to fair weather in most of SEUS, except for the southwest portion of the region. In other words, the fair weather in late fall in SEUS is due to frequent WT4 (MRVR). WT5 (WPT) is more frequent in the first half of the winter (from mid-November to early January) with a contrasting of rainy weather west of the Appalachian and fair weather along the east coast. Finally, WT6 is less frequent before late November and then somewhat evenly distributed after that. Comparing to the warm season WTs with a dominant summer WT (Florida High Zone that occurs about 60% of days in July, see [9]), there is no single dominant WT in the winter.

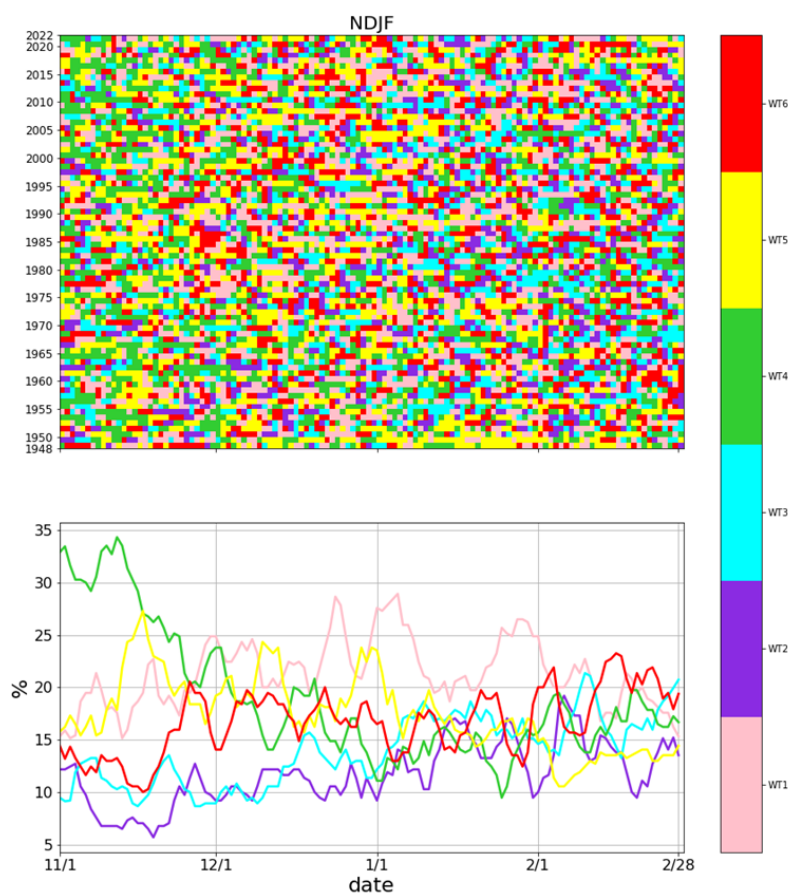


Figure 4. The 74-year (1948–2022) averaged frequency (percentage) of the 6 WTs from November 1 to February 28 (120 days/year). A 5-day running average was conducted. Color: WT1 (pink), WT2 (purple), WT3 (light blue), WT4 (green), WT5 (yellow), WT6 (red).

Figure 5 shows the persistence of the WTs. For all 6 WTs, over 50% of their respective occurrences persist for only one day (and progress to other WTs the next day), and 20% persist for 2 days, 5–10% persist for 3 days, and very few can persist for 4–12 days. These are short compared to the summer WTs that can persist up to 20 days [9]. Winter weather is mostly controlled by eastward propagating baroclinic waves, and therefore, they have short persistence and fast transition/progression from one WT to another.

The progression of the WTs is shown in Fig. 6. WT1 significantly progresses to itself (i.e., persistence, due to its nearly zonal flow) and also likely progresses to WT4 (shown by red bars). Actually, persistence is statistically significant for every WT (red bar). Besides persistence, the preferred progression loop is from WT2 (ECT) to WT3 (OECT), to WT4 (MRVR), to WT5 (WPT), to WT6 (PT), and then back to WT2. The circulation patterns of WTs 2–6 (Fig. 3) indicate that the above progression loop corresponds to the eastward propagation of westerly waves.

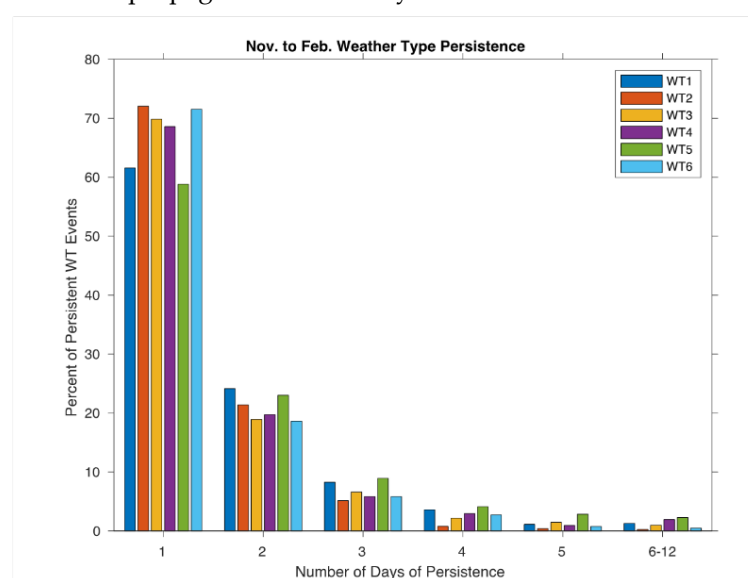


Figure 5. Persistence of the 6 WTs. presented by different color bars, respectively. The x-axis is the length of consecutive days with the same WT. The last group of bars are the sums of persistent events of 6 to 12-days.

NDJF WT Progression (Southeast US)

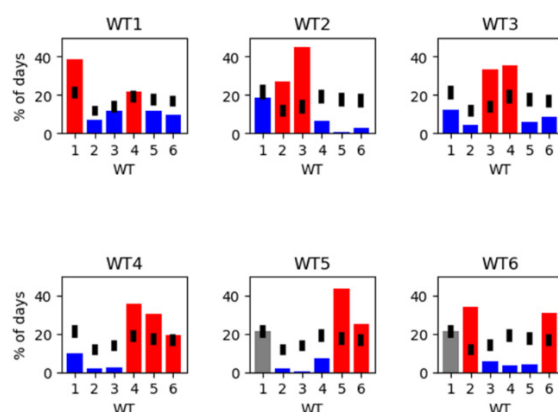


Figure 6. The progression of the WTs. The 95% confidence interval of the WT background frequency (the occurrence of WT transition due to chance) is shown with black shading. Frequencies significantly greater than the background frequency are shown with red bars, frequencies significantly less than the background bars are shown with blue bars.

We also checked the long-term trend in the 1948–2022 time series of the frequency of each of the WTs in November to February (figure not shown) and did not find significant trend based on the Mann-Kendall test at the 95% significance level.

3.3. ENSO impacts on rainfall in the SEUS in the winter

After analyzing the frequency and spatial pattern of the WTs, now we can use them to interpret the ENSO impact on the winter precipitation in the SEUS. As can be seen from Fig. 1, the monthly climatology is similar among each month in deep winter months of December to February (DJF), therefore, we will focus on DJF.

As did by Ohba and Sugimoto in [14] for the ENSO impact on precipitation in Japan, we also quantified the dynamic and thermodynamic contributions of ENSO to the winter precipitation in the SEUS. El Niño and La Niña have opposite impact, but may not be totally asymmetric, so we will check their impact separately. The anomalous precipitation associated with El Niño (denoted by En) can be written as:

$$\delta P = P^{En} - P^{Clim} = \sum_{i=1}^K p_i^{En} f_i^{En} - \sum_{i=1}^K p_i^{Clim} f_i^{Clim} = \delta P_{thermo} + \delta P_{dyn} \quad (1)$$

$K=6$ is the number of the WTs; P^{En} and P^{Clim} are the precipitation in El Niño years and the climatology (all year average), respectively; p_i^{En} and p_i^{Clim} are the precipitation of WT i ($i=1$ to 6) in El Niño years and climatology, respectively; f_i^{En} and f_i^{Clim} are the frequency of WT i in El Niño years and climatology, respectively. δP_{thermo} and δP_{dyn} are the thermodynamic and dynamic contribution to the anomalous precipitation, respectively, as follows:

$$\delta P_{thermo} = \sum_{i=1}^K (p_i^{En} - p_i^{Clim}) \frac{f_i^{En} + f_i^{Clim}}{2} \quad (2)$$

$$\delta P_{dyn} = \sum_{i=1}^K \frac{p_i^{En} + p_i^{Clim}}{2} (f_i^{En} - f_i^{Clim}) \quad (3)$$

The δP_{thermo} in (2) is resulted from the changes in basic (environmental) field, such as temperature and moisture. The δP_{dyn} in (3) is caused by the changes of frequency of the WTs. The thermodynamic and dynamic contribution to precipitation anomalies in La Niña (denoted by Ln) can be calculated similarly, by replacing En by Ln in (1) to (3).

Figure 7 shows the frequencies of the winter WTs in El Niño, La Niña and all years in DJF (f_i^{Clim} , f_i^{En} and f_i^{Ln} , $i = 1, 6$). In El Niño years, the frequencies of WTs 1 and 5 decrease, and those of WTs 2, 3 and 6 increase, and that of WT4 slightly increases.

Every term in (2) and (3) is calculated and shown in Fig. 8. The thermodynamic contribution (equation 2) and dynamic contribution (equation 3) to the anomalous precipitation in El Niño years are shown in the middle and right column of Fig. 8, respectively, and the sum of the thermodynamic and dynamic contributions by each WT is shown in the left column. The total contribution of the six WTs is shown in the bottom row.

The dynamic contribution (equation 3) is approximately the precipitation pattern shown in Fig. 3 multiplied by the frequency difference between El Niño year and all year (red and gray bars) in Fig. 7. The WTs with a shallow trough (WT1) or a trough of more western location (WT5, WPT) are less frequent, producing negative dynamic contribution from WT1 in the coastal plains and from WT5 in the Mississippi River Valley, respectively (Figs. 8c, 8o). The WTs with deep troughs and more eastern location (WTs 2, 3, and 6) are more frequent in El Niño years. Therefore, positive dynamic contribution is in eastern US and southern US in WT2 and WT6, respectively (Figs. 8f, 8r). The precipitation over land is very light in WT3 (Fig. 3), so the dynamic contribution to precipitation from WT3 is also very small (Fig. 8i). The frequency difference of WT4 is small (Fig. 7), thus the dynamic

contribution to anomalous precipitation is small (Fig. 8l). Summing up dynamic contribution from all six WTs, the total dynamic contribution to the anomalous precipitation (Fig. 8u) features a dipolar pattern of a negative anomaly in the Mississippi River valley versus a positive precipitation anomaly in the east coast states.

The thermodynamic contribution is quite different among the WTs. For each WT, the thermodynamic contribution is also quite different from the corresponding dynamic one. The thermodynamic contribution of WT1 is the negative precipitation anomaly west of the Appalachian Mountain. Therefore, the total contribution of WT1 (Fig. 8a) is the southwest-northeast oriented area of negative anomalous precipitation from the Gulf coast to the Northeast U.S. The thermodynamic contribution of WT2, 3, and 4 are small, except for some localized areas in Florida and Texas. The thermodynamic contribution of WT5 is positive anomalous precipitation in the Gulf coast and Mississippi River Valley region (Fig. 8n), with opposite sign to the dynamic contribution (Fig. 8o). The later is stronger than the former so that the total contribution of WT5 is the negative anomalous precipitation extending from northeast Texas to southwest Pennsylvania (Fig. 8m). The thermodynamic contribution of WT6 is strong (Fig. 8q) with positive anomaly in the Gulf and east coast and negative anomaly in the Ohio River Valley. The combination of thermodynamic and dynamic contributions makes the total contribution of WT6 (Fig. 8p) with positive precipitation in all south and east coastal states and negative precipitation anomaly within the Ohio River Valley. The pattern of the total contribution of WT6 (Fig. 8p) is quite similar to the total El Niño impact shown in Fig. 8s, indicating that most of the precipitation anomaly is contributed by WT6. WT2 increased the positive precipitation anomaly in the east coast; and WT1 and WT5 enhanced the negative precipitation anomaly in the Ohio River Valley.

The relative role of the thermodynamic and dynamic contribution to the precipitation anomaly, in terms of percentage, is location dependent, as seen from Figs. 8s, 8t and 8u. For some places, such as Louisiana, the two contributions are opposite in sign. For the east coast, the two contributions both increase precipitation with slightly stronger contribution from the thermodynamic term. In the Gulf coast, the thermodynamic contribution dominates over the dynamic one, especially in the area close to the coast. In the Ohio River valley, both contributions are negative, therefore reinforce each other, but the thermodynamic contribution looks stronger than the dynamic one.

The La Niña impact on the precipitation in the SEUS is shown in Fig. 9. The patterns are quite similar to the El Niño impact shown in Fig. 8, but with opposite signs. However, there are some asymmetries between the El Niño and La Niña impacts. For example, the dynamic contribution of WT1 is smaller (Fig. 9c), the thermodynamic contributions of WT5 seems weaker and less homogeneous (Fig. 9n), and both thermodynamic and dynamic contribution of WT6 is weaker in La Niña than those in El Niño. The area of positive precipitation anomaly in the Ohio River Valley is larger in La Niña (Fig 9s) than that in El Niño (Fig. 8s).

The current result is consistent with that of Nieto Ferreira et al. [15] who found stronger mid-latitude cyclones and more intense precipitation over a large area in the SEUS during El Niño than La Niña and normal years. In contrast, negative rainfall anomalies are found in the east coast and Gulf coast areas in the SEUS in La Niña years.

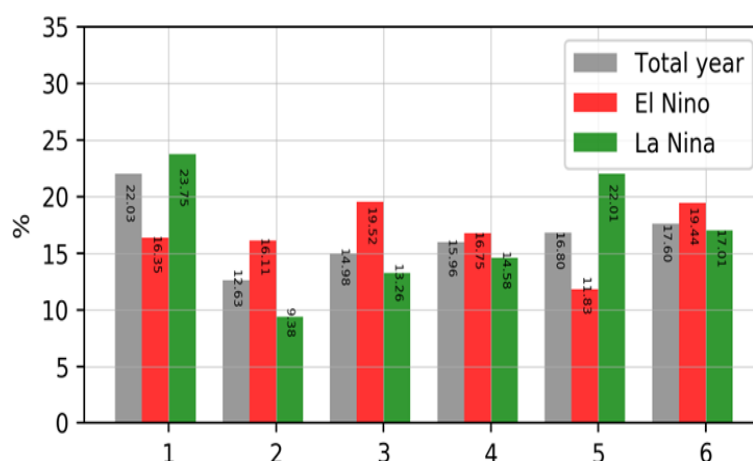


Figure 7. Frequencies of the WTs in all years (grey), El Niño years (red), and La Niña years (green) in December-February of the ENSO years.

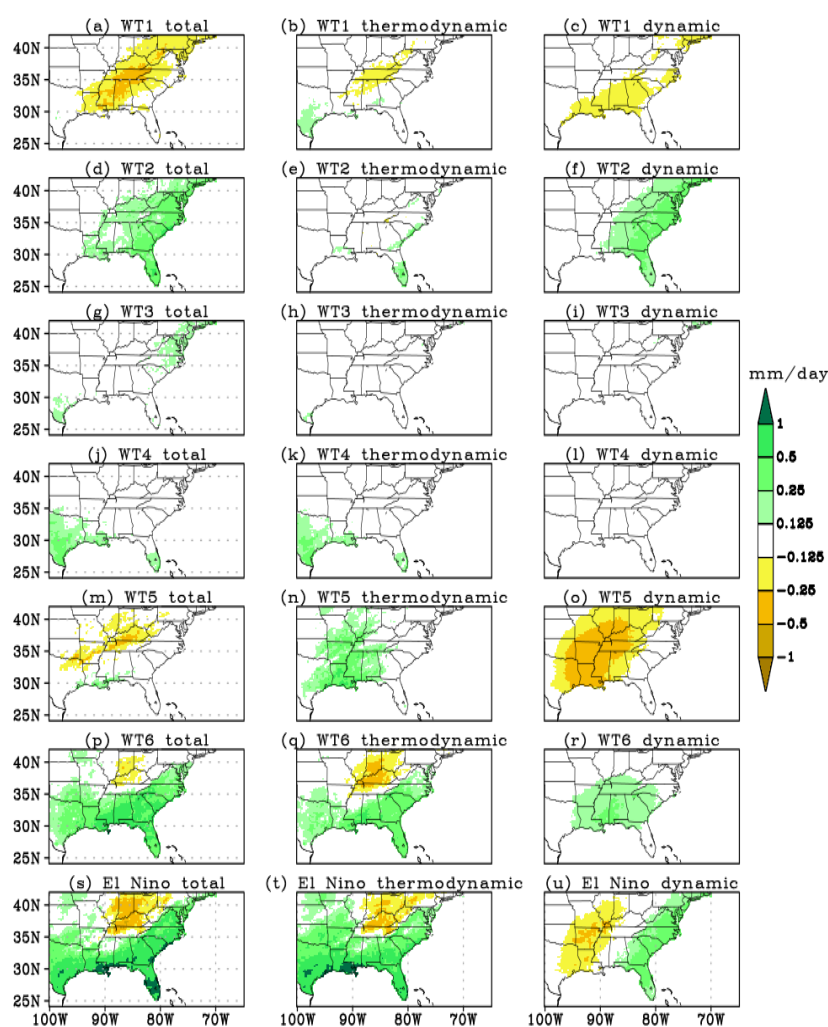


Figure 8. Precipitation anomalies in El Niño years in December-February (panel s). Panels a to r are the contribution of the WTs 1-6 to the anomalous precipitation. Left, middle and right panels are the total, thermodynamic and dynamic contribution from each WT, respectively. Panels t and u are the sum of thermodynamic and dynamic of all WTs, respectively.

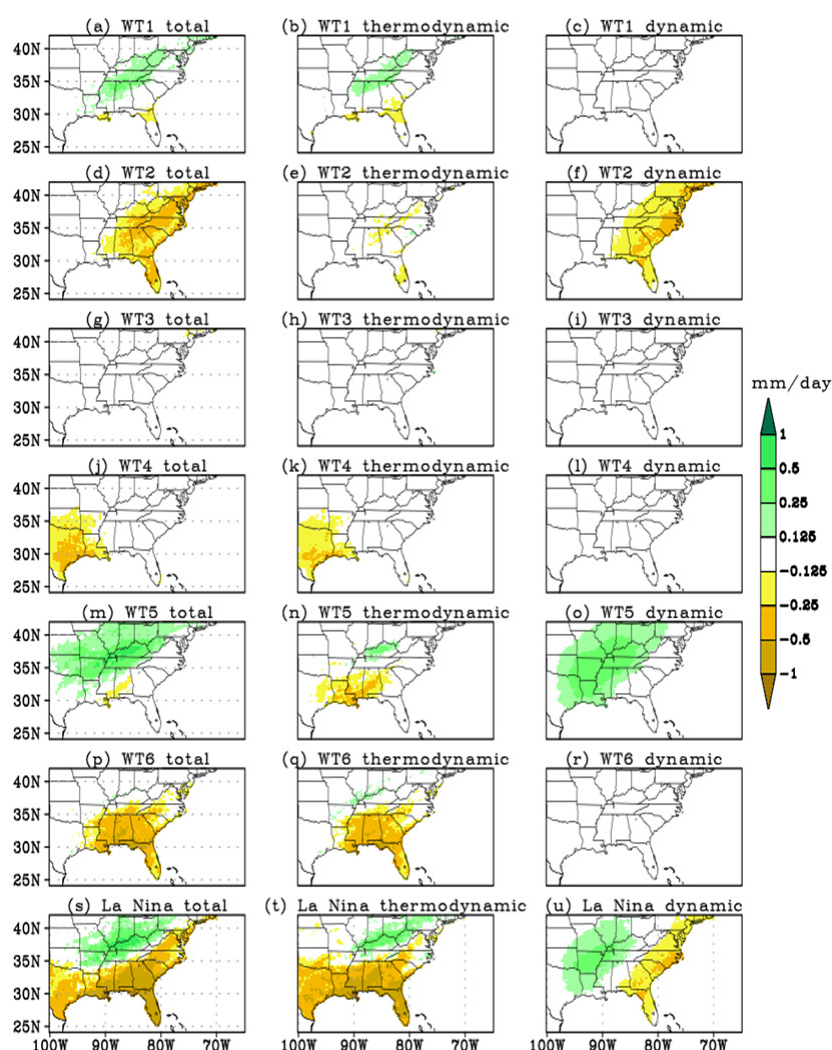


Figure 9. Same as Fig. 8, but for La Niña.

4. Conclusion and Discussion

In order to link weather and climate variability associated with ENSO, we analyzed daily WTs in the SEUS. Six winter weather types have been obtained using k-means clustering analysis of the NRP 850 hPa wind components in November to February of 1948–2022. These six WTs have shorter persistence as compared to the summer WTs in [9]. This is because winter weather is mostly controlled by fast propagating westerly waves. A progression loop comprising WTs 2–6 corresponds to the transient westerly waves.

In our previous study, there are six WTs in the long period of eight months from March to October [9]. Here, there are also six WTs in the short period of four months from November to February. The stronger baroclinicity in the winter generates more fast-moving WTs. In contrast, the quasi-stationary NASH produces long-lasting summer WTs.

The frequency and spatial pattern of the WTs can be effectively used to understand the ENSO impact on winter precipitation in the SEUS. We quantified the ENSO impact in terms of dynamic and thermodynamic contribution to anomalous precipitation, respectively. The dynamic contribution is due to the change of the frequencies of the WTs. There are more frequent WTs of deep east coast trough types and plains trough type (WT2, 3 and 6) and less shallow trough (WT1) and western Plains trough (WT5) in El Niño years. The more frequent east coast trough and plains trough produce above normal precipitation in southeast coastal plains and Gulf coastal plains. The less frequent WT5 tends to reduce rainfall from the Great Lakes southwestward, producing a negative precipitation

anomaly in central U.S. south of the Great Lakes. The thermodynamic contribution is from the change of spatial pattern of environmental variables such as temperature and humidity. In the SEUS, the thermodynamic contribution is as strong as, and even slightly stronger than the dynamic contribution. Particularly, the thermodynamic term of WT6 makes the largest contribution to the anomalous precipitation associated with El Niño. In general, the thermodynamic and dynamic contribution to the anomalous precipitation in La Niña is quite similar to those in El Niño, but with opposite sign. But the influence from WT6 in the La Niña is weaker than that in the El Niño.

Now that we have studied both warm season WTs in March to October [9] and winter WTs in November to February in the current study, respectively, we have obtained a quite complete view of the WTs in the whole year around. Of course, there are some overlapping WTs between warm season and winter WTs, which may be related to the WTs in the transitioning seasons. We have studied the ENSO impact on precipitation from the perspective of the WTs in spring [9] and winter, respectively. It may also be interesting to study the ENSO impacts in the summer and fall using the WTs. Impacts of climate variability at other time scales, such as the intra-seasonal Madden Julian Oscillation and the long-term trend of climate change, also warrant further investigation.

Funding: This work was produced by Battelle Savannah River Alliance, LLC under Contract No. 89303321CEM000080 with the U.S. Department of Energy. Publisher acknowledges the U.S. Government license to provide public access under the DOE Public Access Plan (<http://energy.gov/downloads/doe-public-access-plan>).

Data Availability Statement: The data presented in this study are available on request from the corresponding author. The data are not publicly available due to database access restrictions.

Acknowledgements: We appreciate the constructive comments by the anonymous reviewers which greatly improved the paper.

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